

1 **InSAR Maps Peatlands in the Waikato Region of New Zealand**

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10 **InSAR Maps Peatlands in the Waikato Region of New Zealand**

11 We use remote sensing techniques to assess regional-scale ground deformation in the
12 Waikato Region of New Zealand. Interferometric Synthetic Aperture Radar (InSAR)
13 is applied to measure surface deformation from 2018 to 2022, while digital elevation
14 model (DEM) differencing provides elevation change data from 2013 to 2021. Both
15 methods reveal subsidence within previously mapped peat soil boundaries,
16 confirming ongoing loss of peat soil. InSAR time series show seasonal surface
17 oscillations, aligning with ground-based observations from the same period. The
18 primary driver of subsidence is the artificial drainage of peatlands for agriculture with
19 significant environmental implications for soil conservation, climate change
20 mitigation, water quality, and flood risk management. Our findings demonstrate that
21 satellite-based approaches offer an efficient and scalable method for monitoring
22 peatland subsidence across broad regions, which is not feasible with ground-based
23 methods alone. More moderate subsidence is also observed in low-relief, flood-prone
24 areas beyond the peatlands including the Hauraki Plains, though further validation is
25 needed.

26 Keywords: InSAR, subsidence, LiDAR, satellite, DEM, Waikato, soil, peat.

27 **Introduction**

28 Interferometric Synthetic Aperture Radar (InSAR) is a remote sensing method
29 that utilises satellite-based radar to investigate ground surface deformation. InSAR was
30 previously used in New Zealand to evaluate deformation related to hydrothermal
31 activity and magmatism (Hamling et al., 2016; Hamling, 2020; Harvey, 2021), tectonic
32 processes (Hamling et al., 2022), and slope creep in steep rural terrain, caused by
33 seasonal swelling/shrinking behaviour of clay-rich soils (Cook et al., 2023), and vertical
34 land motion over coastal strips in New Zealand, using Sentinel-1 InSAR and GNSS data
35 (Kearse, 2024).

36 Here we use LiCSBAS an open-source package to carry out InSAR time series
37 analysis using LiCSAR products (July 2018 to July 2022) across 8,000 km² of mostly
38 low relief pastoral land in the Waikato Region of New Zealand. LiSCAR data products

39 are made available by the Centre for Observation and Modelling of Earthquakes,
40 Volcanoes and Tectonics (COMET), and their suitability for measurement of soil
41 deformation in grasslands was previously demonstrated. For example, Xu et al. (2022)
42 measured regional-scale surface deformation of permafrost in the Qinghai–Tibet
43 Plateau, an area covered by alpine meadows, grasslands and deserts, and Orhan et al.
44 (2023) identified widespread (~40,000 km²) subsidence in the grasslands and croplands
45 Konya Basin, Türkiye, resulting from agricultural groundwater usage.

46 We also use DEM differencing, comparing an older satellite-based DEM (2013
47 Copernicus 30m) to a newer LiDAR DEM (2021), to validate LiCSBAS results. DEM
48 differencing was previously used to provide confirmation of InSAR subsidence maps in
49 Taupo, New Zealand (Harvey et al., 2019).

50 Historically, New Zealand was dominated by forest below the alpine treeline,
51 but about 1000 years of Polynesian and European colonisation has resulted in the
52 destruction of nearly three-quarters of the indigenous forest cover (Ewers et al., 2006).
53 As of 2024 approximately 52.5% of the land area of the Waikato Region is covered by
54 high producing exotic grassland (LCDB5, <http://www.lris.scinfo.org.nz>), mostly utilised
55 for dairy farming and livestock grazing.

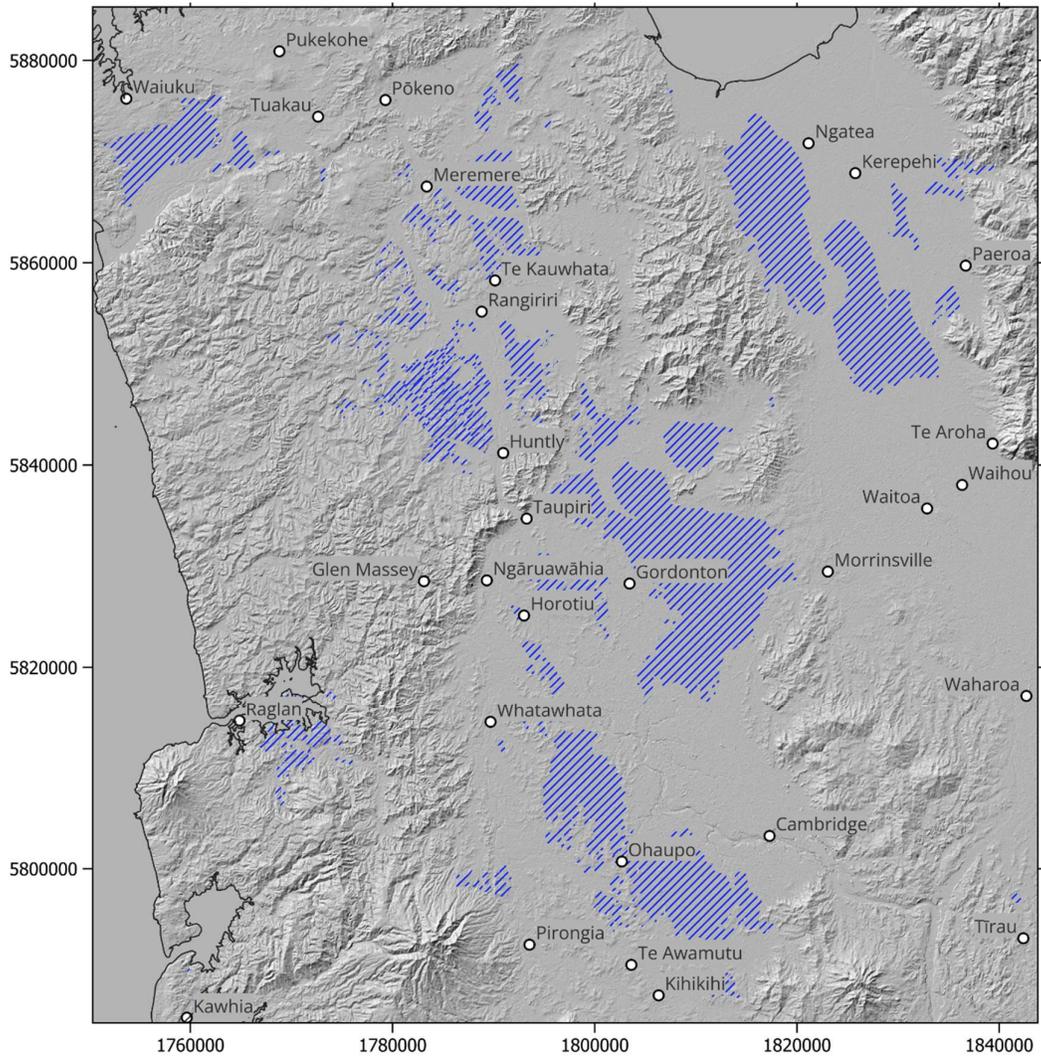
56 New Zealand once had approximately 1660 km² of peat wetlands, around 0.7%
57 of the total land surface area. Of this, approximately 940 km² was in the Waikato
58 Region (*Figure 1*). Over 80% of Waikato peatlands have now been drained for pastoral
59 use (since the early 1900s) (Davoren, 1978; Pronger et al., 2014; Layton, 2022). This
60 has resulted in widespread ongoing subsidence from soil densification and loss of
61 carbon through oxidation of organic matter (19–34 mm year⁻¹) (Schipper and McLeod,
62 2002; Pronger et al., 2014; Glover-Clark, 2020). However, drainage is not limited to

63 peatlands, being widespread throughout low relief areas of Waikato in soils with
64 variable organic carbon content, and related subsidence may be widespread.

65 Drainage benefits agriculture because it allows the soil to drain sooner after
66 heavy rain, increasing grazing days and soil respiration, which improves productivity
67 (Monaghan, 2004). The main environmental impacts of drainage are i) soil shrinkage
68 and subsidence, ii) acceleration of lateral subsurface contaminant flow from land to
69 water (Pearson, 2015), and iii) oxidation of organic matter (soil nutrient loss and
70 increased CO₂ emissions to the atmosphere) (Schipper and McLeod, 2002; Pronger et
71 al., 2014, Glover-Clark, 2020; Layton, 2022). InSAR has been utilised elsewhere to
72 measure surface motion of drained peatlands as these soils are particularly important
73 sources of atmospheric CO₂ (Alshammari et al., 2020; Hoyt et al., 2020; Tampuu et al.,
74 2023; Hrysiewicz et al., 2024).

75 Soil subsidence incurs direct economic costs for farmers; the need to deepen
76 drains continually to match the rate of subsidence further lowers the water table,
77 perpetuating a drainage–subsidence cycle (Pronger et al., 2014). The economic
78 productivity of soils can be lost if they sink below sea level or if the underlying mineral
79 layer is unsuitable for agriculture (Ingebritsen et al., 1999; Pronger et al., 2014).
80 Progressive deepening of drains and streams is also implicated in stream bank erosion,
81 enhancing sediment loss rates in developed catchments (Blann et al. 2009).

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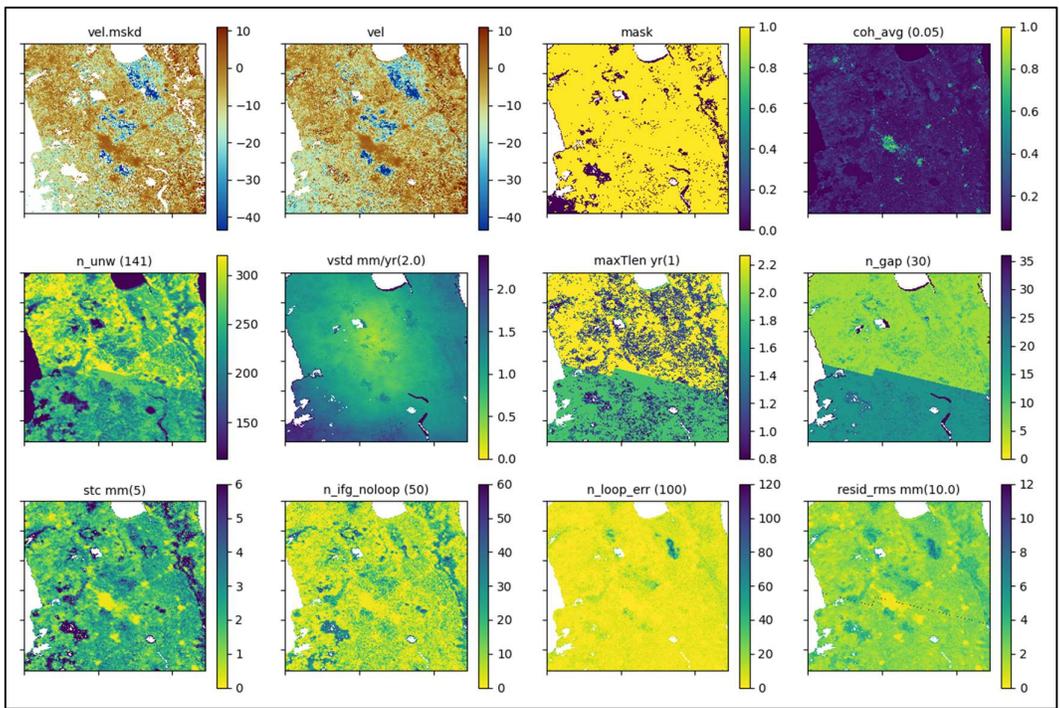
84 Figure 1. Waikato study location overlaid on SRTM DEM (Farr & Kobrick, 2000).
 85 Cross hatching indicates peat soil areas (<https://iris.scinfo.org.nz/layer/119585-s-map-soil-classification-soilorder-aug-2024/>).
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87

88 **Methods**

89 For subsidence mapping and time series plotting we use LiCSBAS, a semi-
90 automated InSAR time series analysis package that interfaces with the automated
91 Sentinel-1 InSAR processor (LiCSAR) (Morishita et al., 2020). LiCSBAS processes
92 publicly available LiCSAR data to undertake InSAR time series processing. An
93 advantage of LiCSBAS is loop closure, a quality control process whereby
94 interferograms with many unwrapping errors are automatically identified and discarded.
95 LiCSBAS was run using masking parameters chosen by trial and error to provide a
96 balance between good coverage in grassland areas, while masking dense forest with low
97 coherence (*Figure 2*).

98



99

100 Figure 2. LiCSBAS masking parameters. Numbers shown in the parentheses next to the
101 titles of each noise index are the applied threshold value. See Morishita et al. (2020) for
102 full explanation of masking thresholds.

103

104 Time series for LOS results were obtained using the LiCSBAS interactive time
105 series viewer that provides both filtered and unfiltered series for any selected pixel. For
106 time series in this report, red cumulative displacement points and curves represent
107 unfiltered results, whilst those shown in blue have a spatial filter of 2 km and a temporal
108 filter of ~40 days applied. Filtering helps to reduce noise and highlight seasonal
109 patterns. The viewer automatically fits sinusoidal curves to the time series to show
110 seasonal cycles.

111 For the Waikato, the LiCSBAS dataset was limited to the descending orbit only
112 (complete ascending data not available for the area/time of interest). Deformation is
113 assumed to result from near-vertical subsidence in peatland areas (Layton, 2022;
114 Pronger et al., 2014; Schipper and McLeod, 2002). Estimation of vertical motion uses
115 the approach of Manzo (2006) (*eq. 2*), simplified for the single line-of-sight (LOS)
116 viewing geometry:

$$117 \quad \quad \quad dz \approx d\text{LOS} / \cos \theta \quad (1)$$

119
120 Where dLOS is LOS motion from the descending pass, and θ is the average
121 elevation angle (radians) for the study area.

122 The period 9th of July 2018 to 30th June 2022 included 111 Sentinel-1
123 acquisitions, with 466 interferograms available from the COMET-LiCS portal for
124 approximately 8,000 km² of the Waikato Region (*Figure 1*). Of these 146 were
125 automatically discarded from further processing by the LiCSBAS loop closure quality
126 control process, leaving a network of 320 interferograms.

127 Small temporal baseline methods including LiCSBAS may be subject to the
128 ‘fading signal’ processing artifact, which can confound results by mimicking and
129 subsidence in cropland environments (Maghsoudi et al., 2022). The uncertainty

130 introduced by this, and the single orbit viewing perspective means deformation rates
131 mapped in this study have unknown precision and may be exaggerated in areas with
132 grass and crops. Despite these uncertainties, LiCSBAS may still permit useful mapping
133 of relative rates, i.e. to determine if elevated subsidence is associated with known peat
134 soil boundaries.

135 An additional deformation map was obtained using an independent method and
136 dataset; DEM differencing was undertaken between i) an Copernicus satellite based
137 DEM from 2013 (average elevation 2011 - 2015 with vertical accuracy $\pm <2\text{m}$ for
138 slopes below 20%) ([https://dataspace.copernicus.eu/explore-data/data-](https://dataspace.copernicus.eu/explore-data/data-collections/copernicus-contributing-missions/collections-description/COP-DEM)
139 [collections/copernicus-contributing-missions/collections-description/COP-DEM](https://dataspace.copernicus.eu/explore-data/data-collections/copernicus-contributing-missions/collections-description/COP-DEM);
140 accessed 01/01/2025), and ii) a LiDAR DEM captured between 5 January and 26 March
141 2021 (vertical accuracy of $\pm 0.2\text{m}$) ([https://data.linz.govt.nz/layer/113203-waikato-lidar-](https://data.linz.govt.nz/layer/113203-waikato-lidar-1m-dem-2021/)
142 [1m-dem-2021/](https://data.linz.govt.nz/layer/113203-waikato-lidar-1m-dem-2021/); accessed 01/01/2025).

143 As for LiCSBAS, DEM differencing also has uncertainties; the vertical precision
144 of the Copernicus DEM ($<2\text{m}$) is large compared to the expected annual rates of
145 deformation ($10\text{'s of mm year}^{-1}$), and the acquisition date (2011 – 2015) means the
146 DEM provides average of elevations in that time period. But just as with LiCSBAS, the
147 approach may still permit useful mapping of relative rates.

148 All reported motions are relative to a reference location on the town of
149 Cambridge that is assumed to be stable (all towns in the Waikato are observed to be
150 stable). Summary statistics (mean and standard deviation) were extracted from the
151 vertical velocity map using QGIS Zonal Statistics. Map grid units are meters (NZTM
152 2000). Map grid units are meters (NZTM 2000).

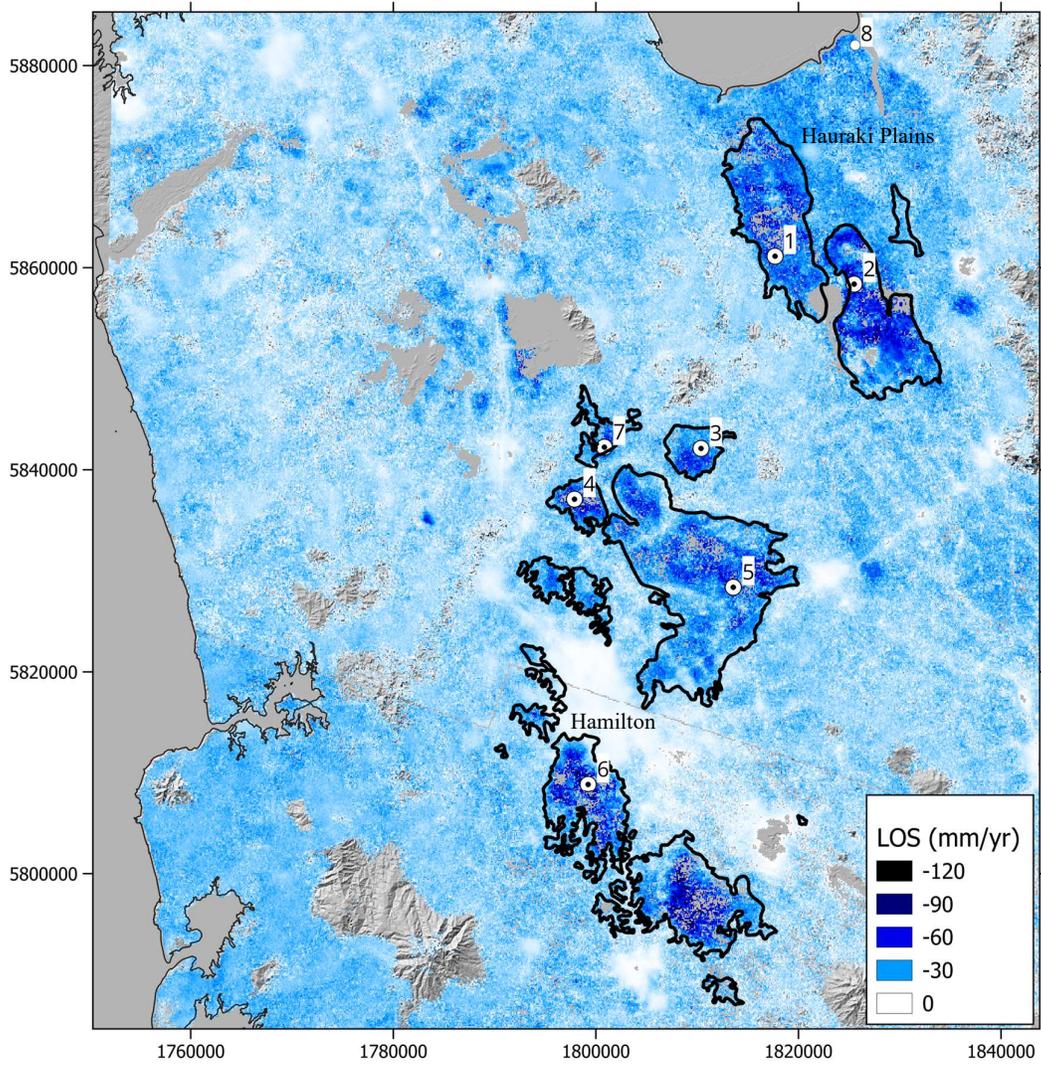
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154 **Results**

155 Mapped InSAR results show strong subsidence within 708 km² of peat soil
156 boundaries (*Figure 3*). Pixels selected from within these boundaries ($n = 6.4 \times 10^4$)
157 averaged -31 mm year^{-1} (LOS) for the 2018 – 2022 period. Assuming the LOS
158 measurements result from vertical subsidence provides an average rate of $\sim -40 \text{ mm}$
159 year^{-1} (*Eq. 1*) for the peatlands, and $\sim -25 \text{ mm year}^{-1}$ for the Hauraki Plains. DEM
160 differencing (2013 Copernicus versus 2021 LiDAR) confirms the subsidence within
161 peat soil boundaries for all southern peatlands, and widespread subsidence in
162 surrounding low-elevation farmland, including the Hauraki Plains (*Figure 4*).

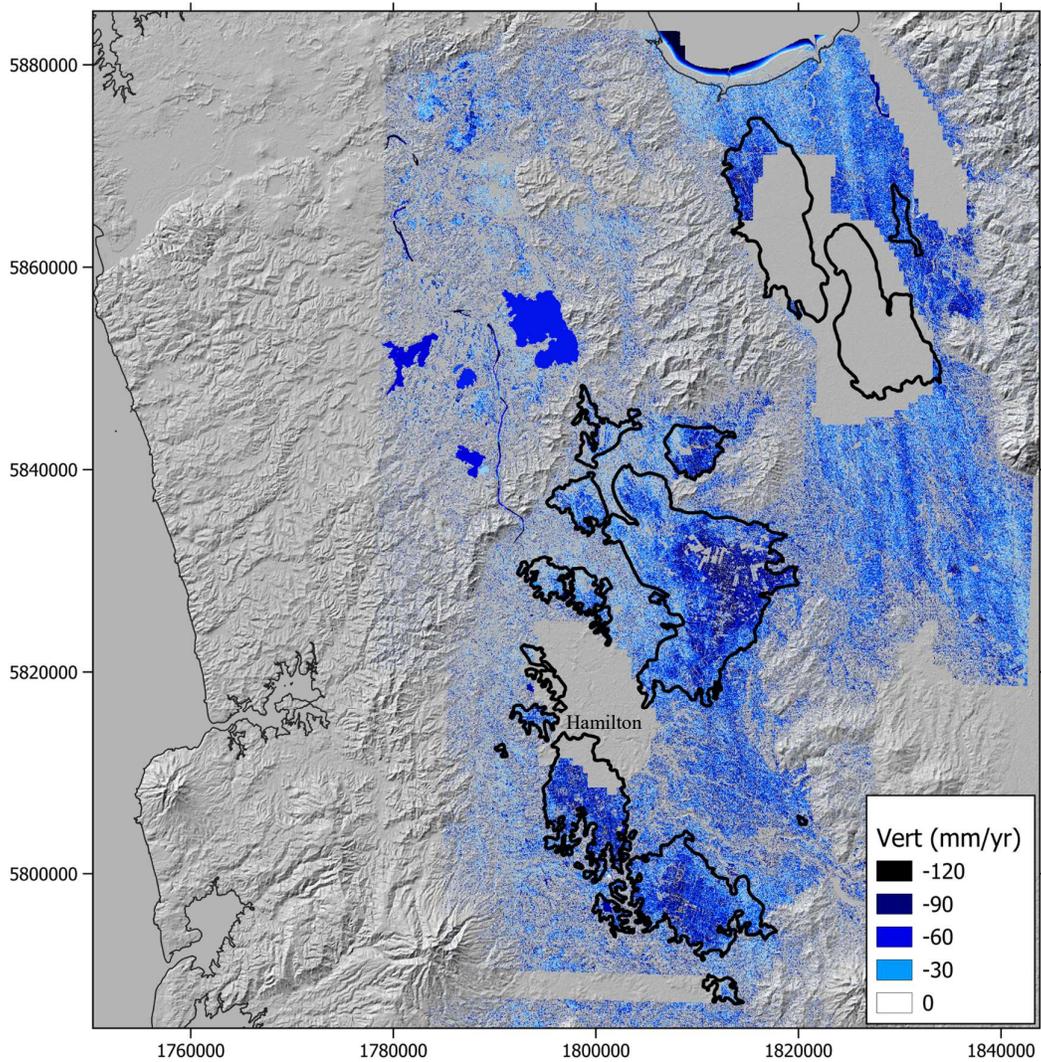
163 Time series for selected peatland locations and non-peat soil on the Hauraki
164 Plains (numbered point locations in *Figure 3*) show well-defined seasonal surface
165 oscillations superimposed on a longer-term subsidence trend (*Figure 5*).

166 Waikato time series were noted to contain many gaps after mid-2020 (see
167 vertical lines in *Figure 5* time series), especially in the southern half of the study area,
168 which derives from a separate Sentinel-1 burst. These gaps divide the series into
169 discrete segments that LiCSBAS aligns by assuming an overall linear displacement
170 trend (Morishita et al., 2020). This process provides good alignment of segments, at
171 least for selected time series in *Figure 5*.



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173 Figure 3. InSAR velocity map showing subsidence associated with peat soils (black
 174 boundaries). Numbers show time series locations (Figure 5).



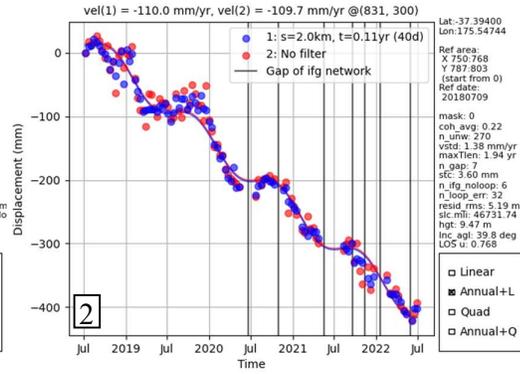
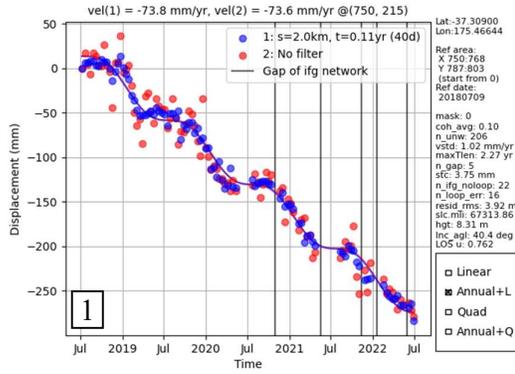
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176 Figure 4. DEM difference map (2021 LiDAR DEM less 2013 Copernicus DEM)
 177 expressed as mean annual motion, showing strongest subsidence associated with peat
 178 soils (black boundaries), which provides independent confirmation of InSAR results
 179 (Figure 3).

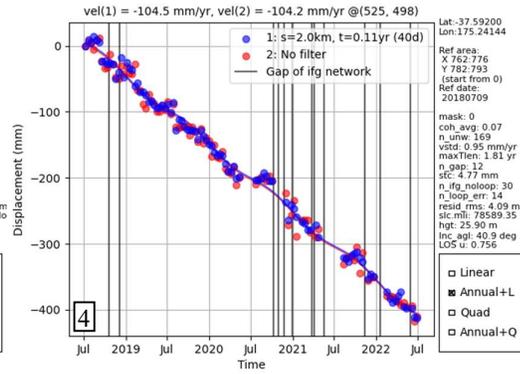
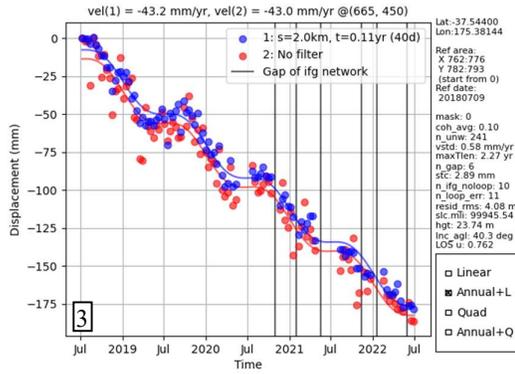
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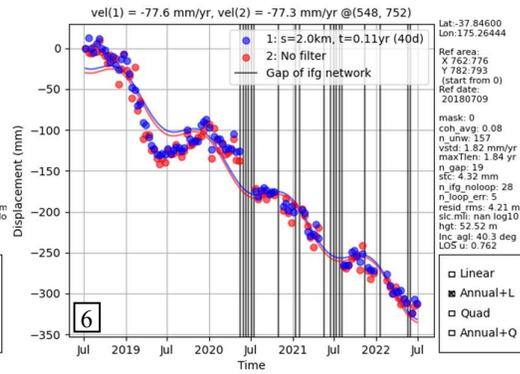
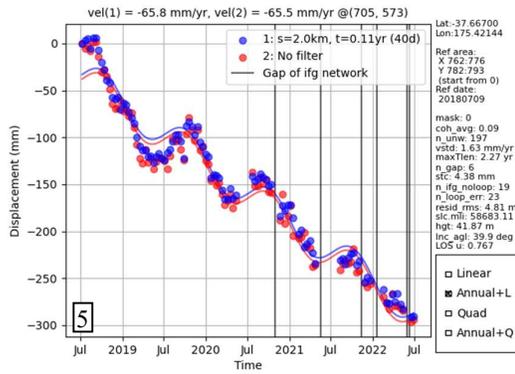
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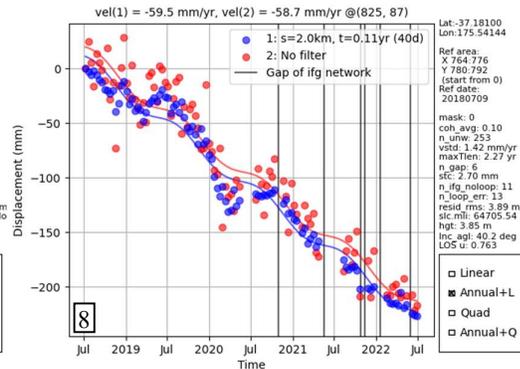
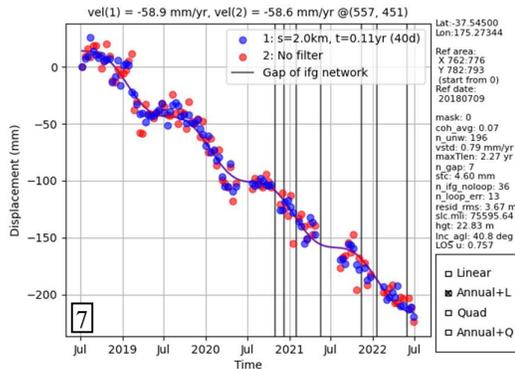


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Figure 5. InSAR time series showing seasonal surface oscillation superimposed on longer term subsidence trend. Numbers correspond to locations in Figure 3.



188 **Discussion**

189 The observation of subsidence in the Waikato peatland soil boundaries mapped
190 by InSAR and DEM differencing is consistent with the findings of previous studies
191 where the causes were reported to be soil densification (shrinkage and consolidation),
192 oxidation and physical compaction resulting from artificial drainage and other
193 agricultural activities (Schipper and McLeod, 2002; Pronger et al., 2014; Glover-Clark,
194 2020; Layton, 2022) (see following section: *Subsidence Mechanisms*).

195 Over 80% of Waikato peatlands have now been drained for pastoral use (since
196 the early 1900s) (Davoren, 1978; Pronger et al., 2014; Layton, 2022). This has resulted
197 in widespread subsidence from soil densification and loss of carbon through oxidation
198 of organic matter, with contemporary estimates of between 19 mm year⁻¹ (Pronger et al.,
199 2014) and 34 mm year⁻¹ (Schipper & McLeod, 2002). The mean rate from LiCSBAS is
200 higher (~ -40 mm year⁻¹) but may be an overestimate because of the ‘fading signal’
201 effect (see *Methods*).

202 Schipper & McLeod (2002) used drained and undisturbed peat soil depths at the
203 Moanatuatua peatland (southernmost peatland in *Figure 3* and *Figure 4*) to estimate
204 mean subsidence (-34 mm year⁻¹), which they attributed partly to densification (63%),
205 and partly to oxidation of organic matter to the atmosphere, i.e. CO₂ degassing (37%)
206 (2.5 to 5.0 ton C ha⁻¹ year⁻¹). A more recent study at Moanatuatua based on repeated
207 GPS measurements (2000 – 2012) provided a lower mean rate (-21 mm year⁻¹) (Pronger
208 et al., 2014). As expected, InSAR provides a much higher mean subsidence (~ -40 mm
209 year⁻¹ vertical) in the area (2018 – 2022), and the DEM differencing estimate is even
210 higher (~ -60 mm year⁻¹) (2013 – 2021). The uncertainty of both LiCSBAS and DEM
211 differencing methods means both mean subsidence estimates must be regarded with

212 caution, but the similarity of deformation maps is compelling, and provides confidence
213 in both methods to delineate peatland areas.

214 Subsidence in the Waikato most concentrated within peatland soil boundaries,
215 but there is also elevated subsidence (~ -25 mm year⁻¹) in soils with lower organic
216 carbon content, i.e. partly mineralised gley soils of the Hauraki Plains near Ngatea and
217 Kerepehi (*Figure 3, Figure 3 and Figure 4*). It is possible this subsidence is an artifact,
218 or exaggerated by the fading signal effect, but interesting to note time-series for both
219 peatland and Hauraki gley soils show seasonal peat surface oscillation (PSO)
220 superimposed on a longer-term subsidence trend (*Figure 5*). If confirmed, subsidence
221 on the low-lying Hauraki Plains could pose a major flood risk, so should be verified by
222 independent methods as soon as possible. CO₂ degassing is another potential issue, as
223 rates for partly mineralised carbon soils (i.e. gleys) are poorly constrained but may be
224 significant (Leiber-Sauheitl et al., 2014; Eickenscheidt et al., 2015; Tiemeyer et al.,
225 2016).

226 Most series show displacement maxima occur after July, which is consistent
227 with recent PSO studies for the same peatlands and period (Glover-Clark, 2020; Layton,
228 2022). These studies measured the oscillation using onsite field methods and correlated
229 shrinking/swelling of soils to seasonal dry/wet conditions. Both studies concluded the
230 oscillation could confound derivation of longer-term subsidence rates (i.e. those based
231 on a limited number of repeat field-based measurements); InSAR provides an advantage
232 in this respect as mean subsidence rates are based on time series of many measurements
233 (~ 12 -day frequency) acquired over multiple years.

234 Surface drainage channels (yellow lines, *Figure 6*) clearly correspond with the
235 subsiding peatland areas, but also the Hauraki Plains. It is interesting to note that the

236 Land Information New Zealand (LINZ) sourced drainage channels may underestimate
237 the true density and extent of drainage channels (Burge et al., 2023).

238 Our results show that InSAR and DEM differencing provides alternate methods
239 for mapping and measuring subsidence of peatlands and other soils at a regional scale.
240 Extrapolation of the InSAR derived subsidence rates indicates about 1m of peat depth
241 will be lost every 25 years. At this rate, large areas of drained peatlands in the Waikato
242 region will be lost much sooner than was previously estimated by Pronger et al. (2014).

243 ***Deformation Mechanisms***

244 In drained peatlands, the primary mechanisms are likely to be soil densification
245 (shrinkage and consolidation) and oxidation. Densification may occur in surface soils
246 where pore volumes are reduced by cycles of drying and wetting, and re-saturation after
247 drying does not restore the soil to its original volume (hysteresis) (Elsaidy,
248 2021). Drainage has resulted in the shrink/swell cycling of peatland soils that were
249 previously saturated throughout the year (Schipper and McLeod, 2002; Pronger et al.,
250 2014; Glover-Clark, 2020; Layton, 2022). To a lesser extent, some soil compaction
251 may occur via a volume-loss process occurring beneath surface soils in the aquifers that
252 underlie low-relief farmland in the Waikato (Griffiths et al., 2023). When groundwater
253 is extracted for irrigation or recharge is reduced by drainage or drought, it causes
254 poroelastic compression of an aquifer's coarse-grained sands and gravel deposits,
255 resulting in subsidence of the overlying surface (Murray & Lohman, 2018). Numerous
256 InSAR studies have correlated subsidence with aquifer compaction caused by
257 abstraction of groundwater for agriculture in arid areas (for example, Castellazzi et al.,
258 2016; Motagh et al., 2017; Navarro-Hernández et al., 2020; Cigna and Tapete, 2021;
259 Peng et al., 2022; Orhan et al., 2023; Lees and Knight, 2024). However, aquifer
260 compression and subsidence in the Waikato may be more likely to result from drainage

261 (reduced recharge to underlying aquifers) than from groundwater abstraction; aquifers
262 lose recharge to surface streams and rivers through drainage (Rissmann et al., 2012).

263 Physical pressure applied by heavy agricultural equipment (e.g. tractors) or high
264 stocking rates is another mechanism that may cause soil densification and contribute to
265 the observed subsidence (Batey, 2009; Keller et al., 2019; Hooijer et al., 2012; Nusantra
266 et al., 2018; Pronger et al., 2014).

267 Oxidation of soil organic carbon (CO₂ release) and associated subsidence is an
268 active area of research in drained Waikato peatlands (Schipper and McLeod, 2002;
269 Pronger et al., 2014; Glover-Clark, 2020; Campbell et al., 2021; Layton, 2022; Pronger
270 et al., 2022) and globally (Dawson et al., 2010; Elsgaard et al., 2012; Hooijer et al.,
271 2012; Hoogland et al., 2012; Tiemeyer et al., 2016; Nusantra et al., 2018; Hoyt et al.,
272 2020; Que et al., 2021; Liang et al., 2024). Drainage lowers the water table, exposing
273 organic material to oxidation, which converts it to carbon dioxide and causes subsidence
274 (Kellner, 2003; Dawson et al., 2010; Hoyt et al., 2020). Soil organic carbon may also
275 be dissolved and mobilised by drainage after rainfall, which provides another
276 subsidence mechanism via mass outflow (Freeman et al., 2004; Guimarães et al. 2013;
277 Rosset et al., 2022).

278 *Stable Areas*

279 The relative stability of towns compared to surrounding pastoral areas may be
280 consistent with the lack of subsidence mechanisms described in the previous section
281 (*Deformation Mechanisms*), or may be an artifact of the fading signal effect (see
282 **Methods**).

283 Assuming the former, towns insulate underlying soils because buildings, roads,
284 and stormwater drainage prevent seasonal shrink/swell cycling from direct exposure to
285 rainfall and sunlight. Agricultural drainage (surface and subsurface) is largely absent in

286 towns, preventing shrinkage from dehydration. Roads are impervious to soil pore fluid
287 exchange, indicating less dehydration from solar heating and less coupling to the
288 atmosphere, which limits oxidation and weathering-related consolidation (Scalenghe
289 and Marsen, 2009).

290 **Environmental Impacts**

291 Soil is the largest terrestrial reservoir of organic carbon, and understanding net
292 inflows/outflows is critical for mitigation efforts (Oertel et al., 2016; Georgiou et al.,
293 2022). Among all soil types, drained peatlands are a particularly concentrated CO₂
294 source (Hoyt et al., 2020; Huang et al., 2021). In New Zealand, drained organic soils,
295 grassland, and cropland alone probably contribute between 7.6% and 9.8% of net
296 carbon emissions (NNE), and InSAR was previously recommended as a method to map
297 subsidence of organic soils (Pronger et al., 2022). The magnitude of these emissions
298 and the uncertainties of current methods mean that it is critical to refine the areal extent
299 of affected soils and associated emission factors. Where oxidation contributes to soil
300 subsidence, InSAR-derived subsidence maps may provide a proxy for organic soils and
301 associated NEE, allowing for improved quantification and mitigation efforts.

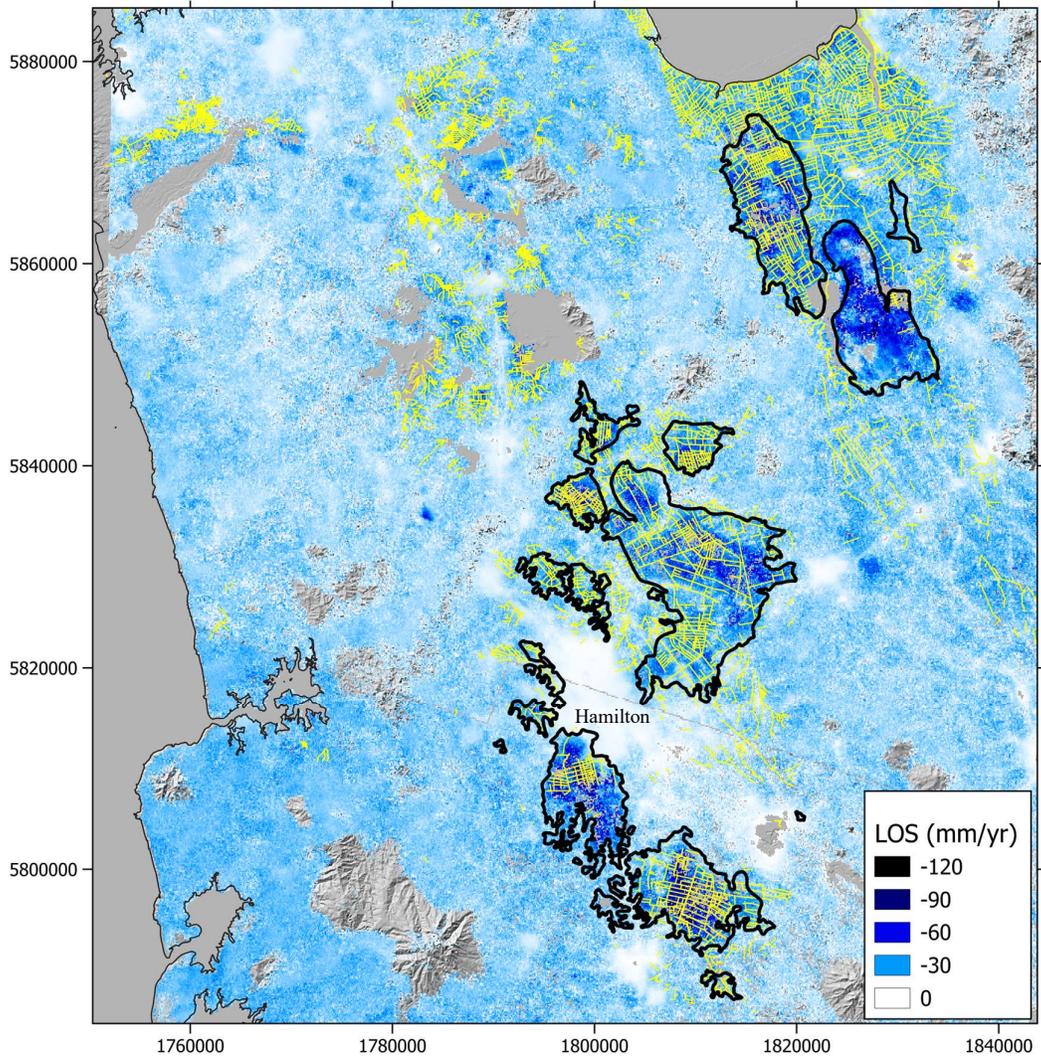
302 The health impacts of nitrate in drinking water are an active area of research in
303 New Zealand (Richards et al., 2022; Rogers et al., 2023). Accordingly, agricultural
304 drainage and associated contamination, especially nitrate from animal urine and
305 mineralisation of highly productive grazing pastures, is a major issue in New Zealand
306 (Monaghan, 2005; Collins et al., 2007; Dench & Morgan, 2021). Both open ditch
307 (surface) and subsurface artificial drainage on farms are widespread in New Zealand
308 (Collins et al., 2007). Such artificial drainage networks are the major conveyance
309 systems for land use derived contaminant loss across large areas of the Waikato.

310 InSAR-derived subsidence maps may provide an indication of contamination source
311 areas through the correlation between drainage and/or irrigation and soil subsidence.

312 Flooding is the most frequent damaging natural hazard in New Zealand (Craig et
313 al., 2021), and climate change is expected to further increase the frequency of extreme
314 rainfall events in New Zealand (Hughes et al., 2021; Paulik et al., 2021). This will be a
315 major issue going forward in low-elevation coastal agricultural areas simultaneously
316 impacted by sea level rise (Craig et al., 2023) and low-relief inland areas subject to
317 inundation from overflowing rivers (Paulik et al., 2021; Griffin et al., 2023). Our
318 results show moderate subsidence in low-lying coastal in the Hauraki District.
319 Confirming subsidence in these areas should be undertaken as soon as possible to allow
320 timely mitigation measures. The situation may be analogous to the Netherlands, where
321 a large part of the western coastal region is covered by organic soils that were
322 historically drained for agriculture; these areas subsided below sea level and now
323 require extensive flood prevention measures (Hoogland et al., 2012).

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Figure 6 InSAR velocity map with peat soils (black boundaries) and agricultural surface drainage channels (thin yellow lines, Land Information New Zealand - <https://data.linz.govt.nz>).

331 **Summary**

332 Here we utilise InSAR analysis (2018-2022), and DEM differencing (2013 –
333 2021) to show distinct areas of subsidence in the northern Waikato Region of New
334 Zealand are neatly contained within known peatland soil boundaries. InSAR time series
335 show the peatlands undergo seasonal elevation changes from rainfall absorption
336 superimposed on a longer-term subsidence trend, which agrees with previous ground-
337 based studies.

338 The primary driver of subsidence is well documented - the artificial drainage of
339 peatlands for agriculture. Our findings demonstrate that satellite-based approaches offer
340 an efficient and scalable method for monitoring peatland subsidence across broad
341 regions, which is not feasible with ground-based methods alone.

342 Widespread subsidence is also mapped across the near-sea-level Hauraki Plains,
343 in both peat and gley soils, but the degree to which this is a processing artifact (fading
344 signal) is unknown. If confirmed by independent methods, this is expected to be a
345 major issue going forward as coastal areas are simultaneously impacted by sea level rise
346 and subsidence; parts of the Hauraki Plains may become unsuitable for traditional
347 agriculture or horticultural use.

348 Future land use decision-making needs to consider the social, economic and
349 environmental consequences of the loss of peatlands, and subsidence of low-elevation
350 farmland elsewhere, which has consequences for soil conservation, climate change
351 mitigation, water quality, and flood risk management. Future work should include
352 extending InSAR coverage to other regions and time periods to determine the full extent
353 of organic soil subsidence in New Zealand. Verification of InSAR results should be
354 undertaken using independent datasets and methods, including repeat LiDAR surveys
355 (point cloud differencing), or conventional ground-based levelling surveys.

356 **Acknowledgments**

357 The authors wish to acknowledge Dr. Yu Morishita for developing the LiCSBAS tool.

358 It is astonishing software, and this work would not have been possible without it. This

359 paper contains modified Copernicus Sentinel data (2018–2022) downloaded from the

360 Centre for Observation and Modelling of Earthquakes, Volcanoes and Tectonics

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592 **Figure Captions**

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594 Figure 1. Waikato study location overlaid on SRTM DEM (Farr & Kobrick, 2000).
595 Cross hatching indicates peat soil areas ([https://iris.scinfo.org.nz/layer/119585-s-
596 map-soil-classification-soilorder-aug-2024/](https://iris.scinfo.org.nz/layer/119585-s-map-soil-classification-soilorder-aug-2024/)).

597

598 Figure 2. LiCSBAS masking parameters. Numbers shown in the parentheses next to the
599 titles of each noise index are the applied threshold value. See Morishita et al.
600 (2020) for full explanation of masking thresholds.

601

602 Figure 3. InSAR velocity map showing subsidence associated with peat soils (black
603 boundaries). Numbers show time series locations (Figure 5).

604

605 Figure 4. DEM difference map (2021 LiDAR DEM less 2013 Copernicus DEM)
606 expressed as mean annual motion, showing strongest subsidence associated with
607 peat soils (black boundaries), which provides independent confirmation of
608 InSAR results (Figure 3).

609

610 Figure 5. InSAR time series showing seasonal surface oscillation superimposed on
611 longer term subsidence trend. Numbers correspond to locations in Figure 3.

612

613 Figure 6. InSAR velocity map with peat soils (black boundaries) and agricultural
614 surface drainage channels (thin yellow lines, Land Information New Zealand -
615 <https://data.linz.govt.nz>).